

Critical loads of nitrogen and sulphur to avert acidification and eutrophication in Europe and China

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Abstract

Introduction Forests and other (semi-)natural ecosystems provide a range of ecosystem services, such as purifying water, stabilizing soils and nutrient cycles, and providing habitats for plants and wildlife. Critical loads are a well-established effects-based approach that has been used for assessing the environmental consequences of air pollution on large regional or national scales.

Materials and methods Typically critical loads of sulphur (S) and nitrogen (N) have been derived separately for characterizing the vulnerability of ecosystems to acidification (by S and N) and eutrophication (by N). In this paper we combine the two approaches and use multiple criteria, such as critical pH and N concentrations in soil solution, to define a single critical load function of N and S.

Results and conclusions The methodology is used to compute and map critical loads of N and S in two regions of comparable size, Europe and China. We also assess the exceedance of those critical loads under globally modelled present and selected future N and S depositions. We also present an analysis, in which the sensitivity of the critical loads and their exceedances to the choice of the chemical criteria is investigated. As pH and N concentration in soil solution are abiotic variables also linked to plant species occurrence, this approach has the potential for deriving critical loads for plant species diversity.

Keywords Critical load function · Exceedance · Nitrogen · Sulphur

Introduction

Critical loads of S and N have been used as an indicators of ecosystem sensitivity to acidification and eutrophication for more than three decades. Critical loads (or sulphate) were first defined in the mid-1980s for surface waters in Canada (e.g. Jeffries and Lam 1993) and the Nordic countries in Europe (e.g. Henriksen et al. 1992). Under the auspices of the Nordic Council of Ministers and the Convention on Long-range Transboundary Air Pollution (LRTAP) of the United Nations Economic Commission for Europe (UNECE) the critical loads methodology was further developed and extended to terrestrial ecosystems (e.g. Nilsson and Grennfelt 1988), resulting in a periodically

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updated manual for the calculation and mapping of critical loads (ICP- Modelling and Mapping 2004). In this context, a critical load is defined as “a quantitative estimate of an exposure to one or more pollutants below which significant harmful effects on specified sensitive elements of the environment do not occur according to present knowledge” (Nilsson and Grennfelt 1988).

Over the last two decades, critical loads have been determined and mapped in several studies for a number of regions: Europe (e.g. De Vries et al. 1994), south-eastern Asia (Hettelingh et al. 1995a), northern Asia (e.g. Bashkin et al. 1995; Reinds et al. 2008), the Arctic region (Forsius et al. 2010), at a global scale (Kuylenstierna et al. 2001; Bouwman et al. 2002) and, for about 15 years, in the P.R. China (Duan et al. 2000; Zhao et al. 2007). In Europe critical loads have been used in integrated assessment (see Reis et al. 2012) to determine cost-effective emission reductions under the UNECE LRTAP Convention (Hettelingh et al. 1995b, 2001, 2007) and, more recently, the air pollution policies of the European Union (Hettelingh et al. 2013). In China, the critical load concept has been applied, e.g., to designate ‘acid rain control zones’ and ‘sulphur dioxide pollution control zones’ (Hao et al. 2001), and there are intentions to use it in future policies on multi-pollutant controls (SO₂, NO_x, particulate matter).

In this paper we focus on critical loads for terrestrial ecosystems. Critical loads are either determined empirically or via models. Empirical critical loads are mostly determined for N, are they derived from field experiments or gradient studies that try to establish a link between deposition and damage to the ecosystem. Recent assessments and summaries of empirical critical loads of N as a nutrient can be found in Bobbink and Hettelingh (2011; for Europe) and in Pardo et al. (2011; for the United States). Alternatively, critical loads are derived from biogeochemical soil models, i.e. models that describe the processes in and chemical state of a soil as a function of its properties and driving forces, such as deposition and harvesting. Whereas in ‘normal use’ models derive the chemical state of the soil as function of the driving forces, critical loads are computed as inverse solutions: given the desired chemical state of the soil, the (maximum) depositions are computed that lead to that state—and these depositions are called critical loads. This task is eased by the fact that critical loads are

calculated for a steady state, i.e. transient processes, such as finite buffers (e.g., of base cations) that can delay the impact of depositions, are neglected—the aim being deposition limits that avoid harmful effects in the long run.

Typically, critical loads have been calculated for effects due to acidity (by S and N) and for eutrophying effects by N (Sverdrup and De Vries 1994; Posch and De Vries 1999; De Vries et al. 2003; ICP Modelling and Mapping 2004), using independent chemical criteria relevant for the respective effect. In this paper we combine the two approaches and use multiple criteria, such as critical pH and N concentrations in soil solution, to define a single critical load function of N and S. As pH and N concentration are abiotic variables also linked to plant species occurrence (e.g., Belyazid et al. 2011), this approach has the potential to derive critical loads for criteria related to plant species diversity.

In the following we first derive the critical load function for N and S and then describe the data bases used to calculate critical loads for forests and semi-natural vegetation in Europe and China. We use these data to compute and map critical loads for both regions, and assess the exceedances of those critical loads for different past and future N and S depositions. In a separate section we conduct an analysis by varying the values of the chemical criteria to investigate the sensitivity of the critical loads and their exceedances to the choice of those criteria. We conclude the paper by a discussion of possible generalisations of the critical load function and the uncertainties in critical load calculations.

Methods and materials

Modelled critical loads of N and S

Critical loads are computed with the Simple (or Steady-state) Mass Balance (SMB) model, which links deposition(s) to a chemical variable (the ‘chemical criterion’) in the soil or soil solution that is associated with undesirable effects (e.g. damage to fine roots) that occur when the ‘critical limit’ of that criterion is violated. Below we shortly summarise the methods for calculating the critical load of nutrient N (eutrophication) and the critical load of acidity, before combining them into the critical load function of N

and S (for details see ICP- Modelling and Mapping 2004).

The critical load (CL) of nutrient N, $CL_{nut}N$ (e.g. in eq/ha/year), is derived from the steady-state mass balance of N in a single-layer soil compartment as:

$$CL_{nut}N = N_i + N_u + N_{de} + Q[N]_{crit} \tag{1}$$

where N_i , N_u and N_{de} are the long-term average immobilisation in soil, net uptake (i.e. removal by harvest) and denitrification fluxes of N, respectively, Q is the percolating water flux leaving the root zone, and $[N]_{crit}$ is the critical concentration of N, above which the ‘harmful effect’ occurs. Equation 1 is the simplest form of the critical loads equation, in which all N fluxes are independent from the N deposition. If denitrification is modelled as proportional to the net input of N, i.e. $N_{de} = f_{de} \cdot (N_{dep} - N_i - N_u)$ with a site-specific constant f_{de} , $CL_{nut}N$ is derived as:

$$CL_{nut}N = N_i + N_u + \frac{Q[N]_{crit}}{1 - f_{de}} \tag{2}$$

The critical load of S acidity is derived from the charge balance in soil solution and obtained as:

$$CL_{max}S = BC_{dep} - Cl_{dep} + BC_w - Bc_u - Q[ANC]_{crit} \tag{3}$$

where $Bc = Ca + Mg + K$ and $BC = Bc + Na$, the indices dep , w and u refer to deposition, weathering and net uptake (i.e. removal by harvest), respectively, and $[ANC]_{crit}$ is the critical concentration of acid neutralising capacity (ANC). $CL_{max}S$ is the maximum deposition of S, S_{dep} , that the ecosystem can stand without causing harm, that is, as long as the N deposition, N_{dep} , is immobilised or taken up, i.e. stays below $CL_{min}N = N_i + N_u$. If $N_{dep} > CL_{min}N$, however, the critical load is a function of both S and N deposition; and for $S_{dep} = 0$ the maximum CL of N is derived as:

$$CL_{max}N = CL_{min}N + CL_{max}S / (1 - f_{de}) \tag{4}$$

The three quantities $CL_{max}S$, $CL_{min}N$ and $CL_{max}N$ define the critical load function of acidifying N and S, depicted in Fig. 1a. For details on the derivation of these equations and their underlying assumptions see ICP- Modelling and Mapping (2004).

The most widely used criteria for avoiding ‘harmful effects’ of acidity are a critical aluminium (Al) to base cation ratio, a critical Al concentration, or a critical pH

(ICP- Modelling and Mapping 2004), as used in this paper. In case of a critical $[H]_{crit} = 10^{-pH_{crit}}$, the critical ANC leaching can be computed as:

$$[ANC]_{crit} = -[H]_{crit} - K_{gibb}[H]_{crit}^3 + \frac{K_1 \cdot DOC}{K_1 + [H]} \tag{5}$$

In Eq. 5 it is assumed that Al-(hydr)oxides are dissolved according to a gibbsite equilibrium; DOC is the concentration of dissolved organic acids, the dissociation of which is modelled by assuming them monoprotic with constant K_1 (ICP- Modelling and Mapping 2004). The bicarbonate concentration has been neglected in Eq. 5.

To date, critical loads of acidity and nutrient N have been treated separately. However, in the case of plant diversity, both the availability of N and the acidity status of the soil are of importance. This suggests treating both types of critical loads in conjunction; and the simplest way is to intersect the CL function of acidity with the CL of nutrient N, as illustrated in Fig. 1b. This CL function is characterised by the four quantities $CL_{max}S$, $CL_{min}N$, $CL_{max}N$ and $CL_{nut}N$ in the N–S-deposition plane (dark grey area in Fig. 1b).

It is important to note that a critical load is a (site-specific) ecosystem property, i.e. it does not depend on the depositions of N or S. A critical load is compared to depositions and if—in the simplest case— $ExN = N_{dep} - CL_{nut}N$ is greater than zero for a given N deposition, the critical load of nutrient N is said to be exceeded, and the ExN is called the exceedance. For critical load functions, the computation of exceedances becomes more involved: the exceedance is the sum of ExN and ExS , which are calculated for a given CL function for the different cases of N_{dep} and S_{dep} depositions according to Fig. 2 (for technical details on the calculations see ICP Modelling and Mapping 2004).

To obtain a single exceedance value for a grid cell (for mapping) or an entire region (for tabulation), the average accumulated exceedance, AAE , has been defined:

$$AAE = \sum_{j=1}^n A_j Ex_j / \sum_{j=1}^n A_j \tag{6}$$

where n is the number of ecosystems in the region (grid cell), Ex_j the exceedance of ecosystem j and A_j the area this ecosystem represents; in other words, the AAE is the area-weighted exceedance of the

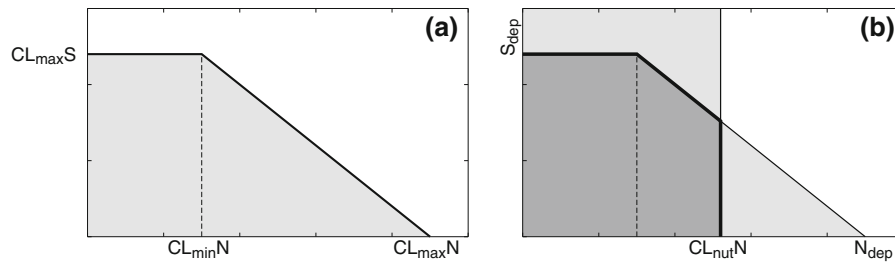


Fig. 1 **a** Critical load function of acidity characterised by the three quantities CL_{maxS} , CL_{minN} and CL_{maxN} . The grey area indicates N and S depositions for which critical loads are not exceeded; **b** acidity CL function intersected with the critical

load of nutrient N, $CL_{nut}N$, defining the N and S CL function. For N and S depositions in the dark grey area neither acidity nor nutrient CLs are exceeded; in other words, neither the acidity nor the nutrient N criterion is violated

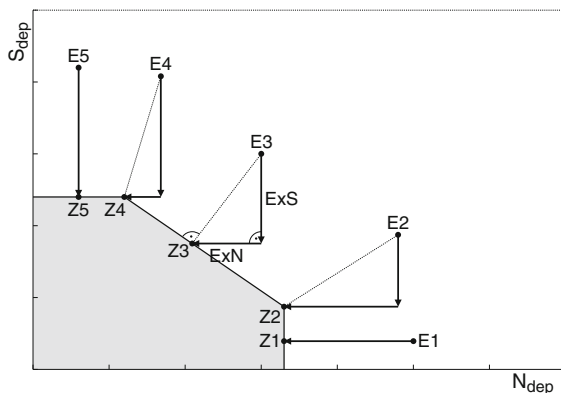


Fig. 2 Illustration of the different cases for calculating the CL exceedance, $Ex = ExN + ExS$, for a given N and S load function for different cases of N and S depositions (E1–E5)

individual CLs. To make the AAE a meaningful quantity, Ex_j has to be set to zero in case of non-exceedance (see Posch et al. 2001).

Input data

The required input data for the calculation of critical loads on a regional scale consist of spatial information on climatic variables, base cation deposition and weathering, net nutrient uptake (removal by harvest) and N transformations. In this study critical loads were calculated for forests (EUNIS code ‘G’; see Davies et al. 2004) and (semi-)natural vegetation (‘D’: mires, bogs and fens; ‘E’: natural grasslands and ‘F’: heathland, scrub and tundra). In the following the data sources and necessary transformations to obtain those inputs are described separately for Europe and China.

Europe

The European area was restricted to land areas west of 32°E, the area for which the most detailed and complete data sources were available.

- (a) *Land cover* We used the European part of the Global Land Cover 2,000 map with 1 km resolution (Bartholome et al. 2002).
- (b) *Soils* The European Soil Database v2 map (JRC 2006) at a scale 1:1 M was used for Europe (including Belarus, Ukraine and the European part of Russia). The soil maps are composed of associations, each map polygon representing one soil association. Every association, in turn, consists of several soil typological units (soil types) that each occupy a known percentage of the soil association, but with unknown location within the association. The soil units on the maps are classified into more than 200 soil types (European Soil Bureau Network 2004), with associated attributes such as soil texture, parent material class and drainage class. Six texture classes (including peat) are defined, based on clay and sand content (FAO-UNESCO 2003). Drainage classes were derived from the dominant annual soil water regime.
- (c) *Meteorology and hydrology* Long-term (1961–1990) average monthly temperature, precipitation and cloudiness were derived from a high resolution European database (New et al. 1999) that contained monthly values for the years 1901–2001 for land-based grid-cells of $10' \times 10'$ (approx. 15×18 km in central Europe). Evapotranspiration and runoff (precipitation surplus, i.e. water leaving the root

zone) were calculated with a sub-model used in the IMAGE global change model (Leemans and Van den Born 1994), following the approach by Prentice et al. (1993); a description of this model can be found in Reinds et al. (2008). The bottom of the root zone was set at 50 cm, except for lithosols which were assumed to have a soil depth of 10 cm only.

- (d) *Base cation deposition and weathering*: Base cation deposition for Europe was taken from simulations with an atmospheric dispersion model for base cations (Van Loon et al. 2005). Weathering of base cations was computed as a function of parent material class and texture class and corrected for temperature, as described in ICP Modelling and Mapping (2004).
- (e) *Net growth uptake (removal by harvest), N immobilization and denitrification*: Since CLs describe a steady-state situation, the net growth uptake of base cations and N by forests was computed from the estimated annual average growth of stems and branches and the element contents of Bc and N in these compartments (Jacobsen et al. 2002). Forest growth and harvest data in EU countries was taken from the European Forest Information Scenario (E-FISCEN) model (Schelhaas et al. 2007). Forest growth for the rest of Europe was derived from the EFI database (Schelhaas et al. 1999) that provides measured growth data for about 250 regions in Europe for various species and age classes. For the small parts of Russia mapped, forest stock data by Alexeyev et al. (2004) were used; these data were also extrapolated to Belarus and the western part of the Ukraine. For (semi-)natural vegetation no net removal of biomass was assumed, and thus the net uptake was set to zero. The denitrification fraction f_{de} , needed to compute the denitrification flux (see above), was estimated as a function of soil drainage status and varies between 0.1 for well drained soils to 0.8 for peaty soils (ICP Modelling and Mapping 2004). The long-term net N immobilization was set at $1 \text{ kg N ha}^{-1} \text{ year}^{-1}$, which is at the upper end of the estimated annual accumulation rates for the build-up of stable C–N compounds in soils (ICP Modelling and Mapping 2004).

- (f) Overlaying these maps and merging polygons with common soil, vegetation and region characteristics within grid cells of $0.05^\circ \times 0.05^\circ$ resulted in about 3.17 million computational units ('sites') with a total area of 2.97 M km^2 . The determination of critical loads was limited to units larger than 0.5 km^2 , reducing the total number to 1.37 M, covering $\sim 88 \%$ of the study area.

China

- (a) *Vegetation* The Chinese Vegetation Map at a scale of 1:2 M was used. The vegetation units on the maps are classified into 95 types (Wu 1980).
- (b) *Soils* The Chinese Soil Map at a scale of 1:2 M was used. The soil units on the maps are classified into 43 soil types (Xiong and Li 1987).
- (c) *Runoff* The Chinese Runoff Map at a scale of 1:4 M was used (SinoMaps 1984).
- (d) *Base cation deposition and weathering*: Atmospheric emissions of Ca and Mg (K and Na are not included due to their negligible shares of base cation emissions) from major anthropogenic sources in China were estimated based on the emission rate and chemical composition of particulate matter (PM). Zhao et al. (2011) estimated that overall emissions were approximately 5.97 and $0.24 \text{ Tg year}^{-1}$ for Ca and Mg, respectively, in the year 2005. These estimates were much higher than those for Europe of 0.75 – $0.80 \text{ Tg year}^{-1}$ in 1990 (Lee and Pacyna 1999), highlighting the limits of applying Europe's experience to countries like China, where base cations may have a much larger role in atmospheric chemistry and ecological effects. In addition to the anthropogenic emissions, base cation emissions from wind-blown dust particles from arid and semi-arid regions in northern China were roughly estimated to about $2.08 \text{ Tg year}^{-1}$ for Ca and $0.68 \text{ Tg year}^{-1}$ for Mg (Zhu et al. 2004). Using a multi-layer, dynamic Eulerian model (Duan et al. 2007; Zhao et al. 2011), the long-range transport and deposition of base cations was estimated. Results show that the total deposition of Ca in China was about

5.29 Tg year⁻¹ (with 1.54 Tg year⁻¹ as wet deposition and 3.84 Tg year⁻¹ as dry deposition), of which anthropogenic sources contributed 67.1 % (92.4 % of wet deposition and 57.6 % of dry deposition). Samples of all major Chinese soil types have been collected in recent years and analysed for physicochemical properties and mineralogy. These data have been used to compute weathering rates of these soil types with the PROFILE model (Duan et al. 2002).

- (e) *Nutrient uptake and N immobilization* A literature review on average growth rates and element contents of major plant communities in China was carried out by Duan et al. (2004). Based on this dataset, net uptake rates of N and base cations was estimated for each major vegetation type in China and the results were used in the critical load calculations. N immobilization of major soil groups in China was estimated as the total amount of soil N divided by the period of soil formation (Hao et al. 2003).
- (f) *Other parameters* Where needed, these were derived on the basis of ICP- Modelling and Mapping (2004).

Results

Critical loads of N and S

Using the methods and databases described above we calculated the critical load quantities defining the critical load functions of N and S for all sites in Europe and China. Since there are no widely-accepted unique chemical criteria for pH and N, we selected a critical pH of 4.4 (for non-calcareous soils) and a critical N concentration of 1 mg N L⁻¹, independent of location. These values are at the more stringent end of the range of proposed criteria in the literature (Sverdrup and De Vries 1994; Aherne et al. 2001; De Vries et al. 2010). For calcareous soils $CL_{max}S$ was set to 10,000 eq ha⁻¹ year⁻¹, reflecting the very high BC weathering rates (and thus pH) in such soils. In this study, those chemical criteria were applied to both Europe and China, leading to a consistent set of critical loads across both regions.

There are many sites and thus critical loads within a single mapping unit (0.5° × 0.5° grid cell), and rather than displaying the mean, we compute the 5-th

percentile of the distribution of the critical load values within every grid cell, considering that it is the low critical loads, i.e. the most vulnerable ecosystems, that need attention. In Fig. 3, maps of the 5-th percentiles of the maximum CL of sulphur, $CL_{max}S$, and the CL of nutrient N, $CL_{nut}N$, are displayed. The latter is shown here, since for the chemical criteria chosen $CL_{nut}N < CL_{max}N$ in 99.97 % of the sites. In the exceedance calculations the N and S CL function is used to obtain a single exceedance value (see below).

Very low values of $CL_{max}S$ are found in Sweden and Finland and, to a lesser extent, in the Baltic states; values between 400 and 1,000 eq ha⁻¹ year⁻¹ are dominant in Germany, Poland and parts of south-eastern Europe as well as in central-western Spain, i.e. in drier areas (small precipitation surplus). High values dominate in the South and on the west coast of Europe, where they are due to the high precipitation—the main reason why Norway and Denmark differ from the Sweden and Finland (Fig. 3a). The generally increasing north–south gradient in $CL_{max}S$ is due to the increasing base cation deposition (and net uptake), which by and large, increase when going south. Critical loads of nutrient N are much lower than the maximum CLs of S (Fig. 3). These values are even more influenced by the runoff, and thus the lowest values correspond strongly with the areas of low precipitation (surplus) in Europe (Fig. 3b); and the highest values are found on the west coast of Europe and the Alps, regions with high precipitation.

Values of $CL_{max}S$ in north-central and north-western China generally exceed 1,500 eq ha⁻¹ year⁻¹, attributable mainly to high weathering rates and base cation deposition (Fig. 4a). In the northeast and south, however, the values can be lower than 200 eq ha⁻¹ year⁻¹, because of low weathering rates and high base cation uptake. Low $CL_{nut}N$ is commonly found in the north-west, due to very low precipitation and low vegetation N net uptake (Fig. 4b). The values are relatively higher in the south and east, resulting from much higher precipitation and higher soil denitrification rates.

Exceedances

Critical loads provide insight into the spatial distribution of the sensitivity of ecosystems to N and S depositions. To find out whether there is a danger of ‘harmful effects’, critical loads are compared with present and potential future depositions by calculating

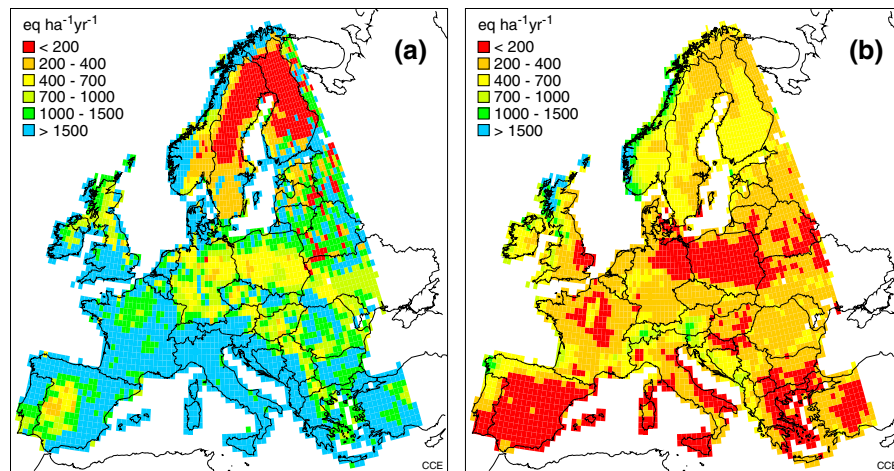


Fig. 3 Gridded ($0.5^\circ \times 0.5^\circ$) 5-th percentile of the critical loads CL_{maxS} (left) and CL_{mutN} (right) for forests and (semi)-natural vegetation in Europe (west of $32^\circ E$)

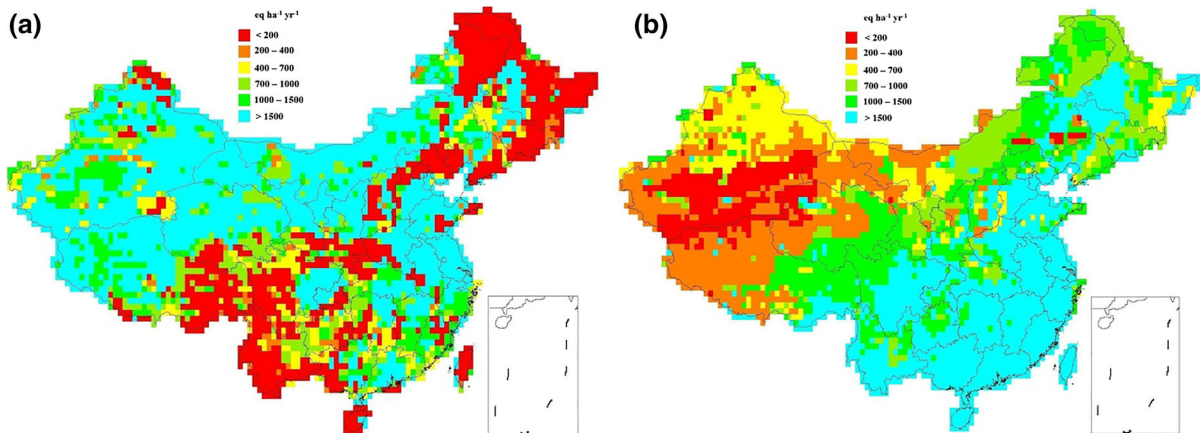


Fig. 4 Gridded 5-th percentile of the critical loads CL_{maxS} (left) and CL_{mutN} (right) in China

the critical load exceedance, i.e. a single number per CL function, as illustrated in Fig. 2. To be able to map exceedances we computed the average accumulated exceedance, *AAE*, as defined in Eq. 6.

As a common and documented basis for exceedance calculations in China and Europe we use S and N depositions available from the Atmospheric Chemistry and Climate Model Intercomparison Project (ACCMIP) (Lamarque et al. 2013). This multi-model deposition dataset provides state-of-the-science, consistent and evaluated gridded ($0.5^\circ \times 0.5^\circ$) global deposition fields (spanning 1850–2100). For the post-2000 atmospheric chemistry, related emissions and concentrations the representative concentration pathways (RCPs) were used (Van Vuuren et al. 2011).

These emission pathways (‘scenarios’) were prepared for the latest assessment of the intergovernmental panel on climate change (IPCC), and thus are more climate-centred rather than looking at N and S emission reduction policies.

Here, in addition to using the 2000 (‘present’) S and total inorganic N ($= NO_y + NH_x$) depositions, we also computed exceedances with depositions for the year 2100 under the RCP4.5 scenario. The RCP4.5 scenario aims at a stabilization at $4.5 W m^{-2}$ (~ 650 ppm CO_2 eq) radiative forcing (see Van Vuuren et al. 2011 for details) and it leads to, by and large, the lowest emissions of S and N in Europe and Asia of the RCP scenario family (see supplementary information of Lamarque et al. 2013). Maps of the

exceedances (AAE) on the $0.5^\circ \times 0.5^\circ$ grid for the years 2000 and 2100 are shown in Fig. 5 (for Europe) and Fig. 6 (for China).

In Europe critical loads in 54.2 % of the ecosystem area are exceeded under the 2000 N and S deposition. Exceedances of critical loads are highest in north-central Europe (Germany, Poland), but also large areas in southern Europe show considerable exceedance. Despite the lowest overall critical loads in Fennoscandia (Fig. 3), large parts are not (or no longer, given that depositions were much higher in the 1970s and 1980s) exceeded, or the exceedance is low. It is interesting to note that computing the exceedances for acidity CLs alone leads to an exceeded area of only 4.7 %,

showing that the N-related part of the CL function is much more stringent. Under the RCP4.5 scenario N and S depositions for the year 2100, exceedances decline essentially everywhere in Europe, both in amount (below $700 \text{ eq ha}^{-1} \text{ year}^{-1}$ almost everywhere) and in area (to about 28 %). And a separate calculation shows that the critical loads of acidity would be virtually non-exceeded in all of Europe (exceeded area $<0.1 \%$).

In China, critical loads in about 50 % of the ecosystem area are exceeded under the 2000 N and S deposition (Fig. 6a). Exceedances of critical loads are highest in eastern and south-central China. Those regions have both high S and N depositions, due to the

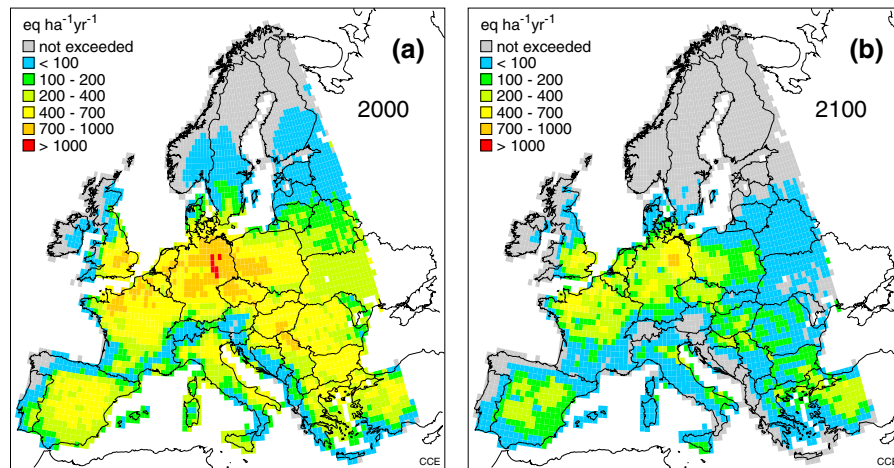


Fig. 5 Exceedance (AAE; see Eq. 6) of critical loads of N and S in Europe in 2000 (left) and in 2100 (right) under the RCP4.5 scenario

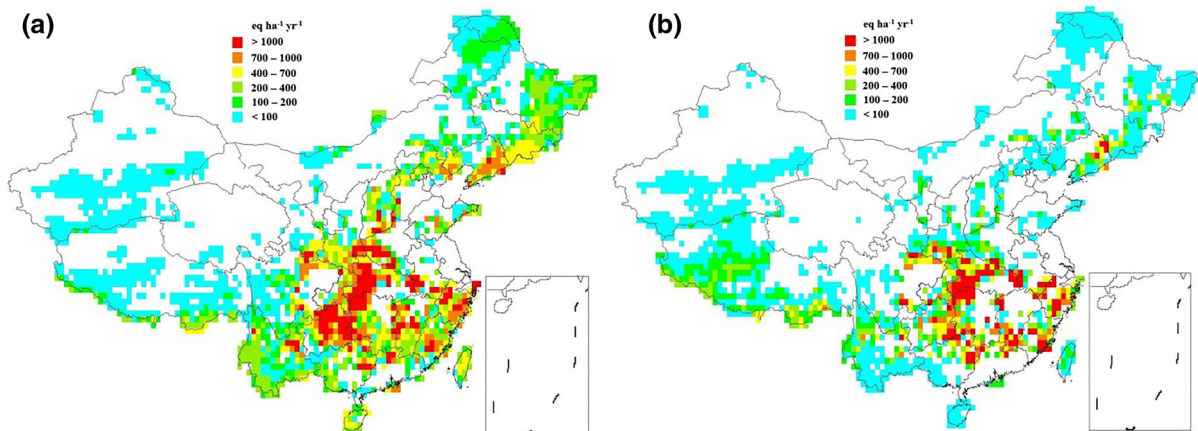


Fig. 6 Exceedance (AAE; see Eq. 6) of critical loads of N and S in China in 2000 (left) and in 2100 (right) under the RCP4.5 scenario

high density of industry, but mainly S exceedance alone, showing that—in contrast to Europe—the S part of the CL function is much more stringent in these areas. However, the critical load exceedance for almost all regions in western China was due to N deposition; but this exceedance was low. Under the RCP4.5 scenario N and S depositions for the year 2100, exceedances decline somewhat in China, mainly visible in the south-eastern part of the country (Fig. 6b). Since S deposition decreases considerably from 2000 to 2100 but N deposition remains about the same, critical load exceedance in 2100 would be mostly caused by N deposition.

Sensitivity to chemical criteria

The uncertainty and sensitivity in critical load calculations on a large regional scale have been studied before (Reinds et al. 2008, Reinds and De Vries 2010). To investigate the sensitivity of the critical loads and their exceedances on the chemical criteria, we computed the $CL_{max}S$ in Europe also for $pH_{crit} = 4.2$ and 4.6. In Fig. 7a the cumulative distribution functions (cdfs) of the $CL_{max}S$ values for the different pH-criteria are displayed. Surprisingly, for small critical loads (i.e. sensitive ecosystems), the $CL_{max}S$ values do not change much for the different criteria, whereas for high acidity critical loads there is a substantial change in the distributions. An analogous analysis was carried out for $CL_{nut}N$ and $[N]_{crit}$. Figure 7b shows, in addition to the base case, the cdfs of $CL_{nut}N$ for $[N]_{crit} = 0.5$ and 2.0 mg N L⁻¹. Obviously, there is a considerable influence, but less than the halving or

doubling of the critical value might suggest. These graphs show that it is important to carefully select the critical values. However, one has to keep in mind that these are extreme results: in practice it will hardly ever be the case that the uncertainties in a critical value will all have to be corrected in the same direction.

A similar analysis was carried out for China. As shown in Fig. 8, CLs are much less sensitive to the changes in the criteria in China than in Europe. In contrast to Europe, the $CL_{max}S$ values change more for small values. As in Europe, the bigger change in the distributions occurs for high nutrient N critical loads. Note that, in China there is a considerable area with $CL_{max}S = 0$; this happens in forest areas where the removal (vegetation uptake and leaching) is higher than the input of base cations (weathering and deposition).

A change in critical loads due to a change in criteria (or any other input parameter) is one thing; however, in applications it is of more relevance how exceedances change under different criteria for the CL function. Using the variation in the criteria as above, the cdfs of the resulting exceedances (AAE) in the year 2000 in Europe and China are displayed in Fig. 9. The results are interesting, reflecting the different pollution climates in Europe and China. For Europe, the change in the critical pH hardly matters. This is due to the fact, that S depositions are already quite low in most of Europe, and thus exceedances are driven by N deposition (alone). This is also reflected in the graphs: The changes in the N criterion (halving and doubling of the $[N]_{crit}$) results in a reduction/increase of the non-exceeded area of about 15 % (around a reference value

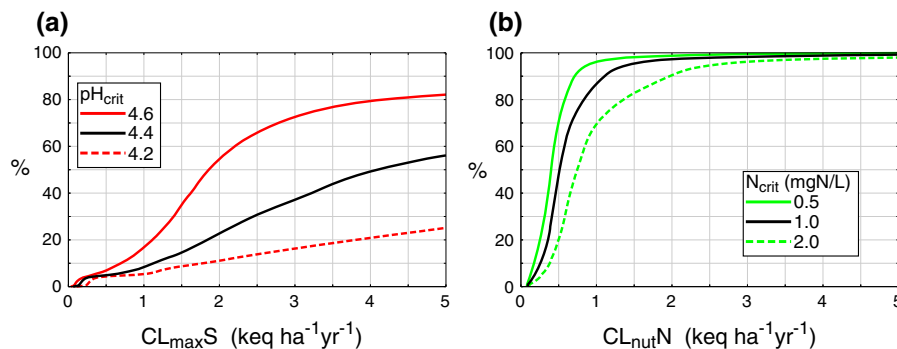


Fig. 7 Cumulative distribution functions (cdfs) of $CL_{max}S$ (a) and $CL_{nut}N$ (b) in Europe for different critical values. The black curves show the base cases ($pH_{crit} = 4.4$,

$[N]_{crit} = 1$ mg N L⁻¹), whereas the red and green curves show the cdfs of $CL_{max}S$ and $CL_{nut}N$, respectively, for the values used in the sensitivity analysis (see legends)

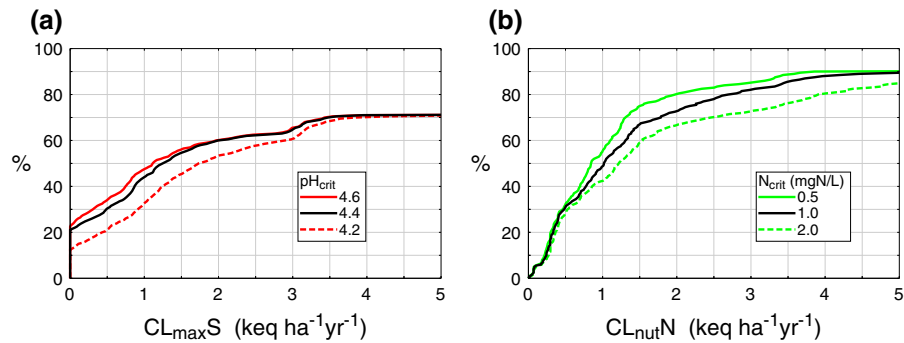


Fig. 8 Cumulative distribution functions (cdfs) of CL_{maxS} (a) and CL_{nutN} (b) in China for different critical values. The black curves show the base cases ($pH_{crit} = 4.4$,

$[N]_{crit} = 1 \text{ mg N L}^{-1}$), whereas the red and green curves show the cdfs of CL_{maxS} and CL_{nutN} , respectively, for the values used in the sensitivity analysis (see legends)

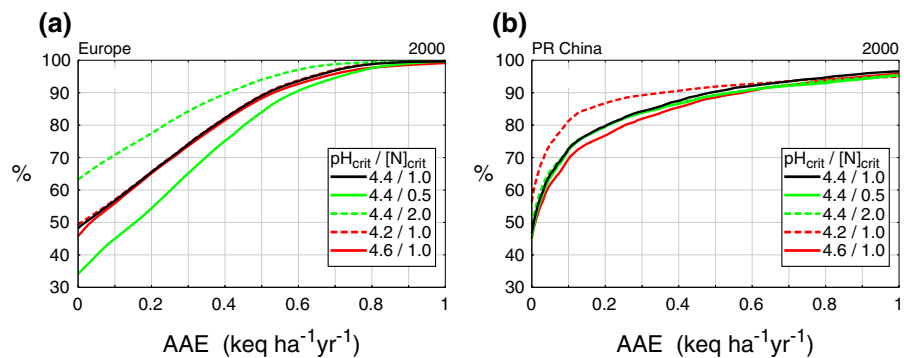


Fig. 9 Cumulative distribution functions of the exceedance (AAE; see Eq. 6) of critical loads in a Europe and b China in the year 2000 for the different combinations of chemical criteria

of 50 %; see intersection of the cdfs with the left axis). In contrast to Europe, the change in the N criterion hardly matters in China. This is due to the fact, that S depositions are main driver of exceedances in China. The changes in critical pH result in a reduction/increase of the non-exceeded area of about 5 % (around a reference value of 50 %).

Discussion and concluding remarks

Combining databases on soil, land cover, climate and forest growth provided detailed maps with more than 3 million receptors for Europe and about 4 million receptors in China, suitable for spatially highly disaggregated critical load calculations, demonstrating that the critical loads methodology can be applied on a large regional scale in different parts of the globe yielding results for comparable assessments.

(pH_{crit} and $[N]_{crit}$). The black line represents the base case ($pH_{crit} = 4.4$, $[N]_{crit} = 1 \text{ mg N L}^{-1}$) and the coloured lines correspond to the values shown in Figs. 7 and 8

In applications to date, critical loads of acidity and eutrophication have been treated separately. However, the growing emphasis on biodiversity (especially plant diversity) renders the separation between acidifying and eutrophying effects inappropriate (especially with respect to N). Therefore we combined in this paper the two approaches into a common critical load function for N and S and computed them on large regional scales (China and Europe). In this exercise we have chosen pH and N concentration criteria without taking into account possible interdependencies between the two variables. However, when looking at the occurrence of plant species or species compositions, the influential abiotic factors, such as pH or [N], are not independent from each other (e.g. Wamelink et al. 2011). Vegetation models that link species occurrence with abiotic factors could be used to derive combined pH-N-limits for the derivation of critical load functions for plant diversity (for a discussion of the issues

involved when dealing with combined criteria of [N] and pH see Posch et al. (2011)). The strong dependence of the critical loads on precipitation shows the limitations of assuming a relationship between N concentration and plant species sensitivity, whereas N availability, or other less fluctuating N quantities, may be more appropriate (De Vries et al. 2010). It should also be mentioned that some authors link the *cumulative* deposition of N to changes in plant species diversity (e.g. Duprè et al. 2010), which would call for a re-formulation of the ‘critical load’ concept.

The overall uncertainty of the critical loads is determined by model uncertainty and parameter uncertainty (see, e.g., Zak and Beven 1999; Skeffington et al. 2007). Parameter uncertainty can be substantial and strongly contribute to uncertainty in critical loads (De Vries et al. 1994). Others have shown that different methods of estimating weathering rates can yield strongly varying results (Hodson and Langan 1999). Model structure is another important source of uncertainty. The SMB model is by design a very simple model of reality; and the simple denitrification sub-model and the assumption of a homogeneous soil layer contribute to uncertainty. Also, previous studies have shown that there is a strong effect of the depth at which the chemical criterion should be met (De Vries et al. 1994).

Although the future N and S deposition scenario used in this paper is motivated by climate change goals, it leads to a strong reduction of present S deposition in Europe and Asia (70–80 %) by 2100. Also N deposition is reduced by about 50 % in Europe, but only by a small fraction in Asia. But as the exceedance calculations carried out in this paper show, only further reductions of (mostly N) emissions will achieve the long-term objective of non-exceedance of critical loads. We think that studies like this could also complement the work carried out by the Task Force on hemispheric transport of air pollution (HTAP) under the LRTAP Convention, which aims at improving the understanding of the intercontinental transport of air pollutants across the Northern Hemisphere.

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